Distribution of Organic Carbon in Surface Sediments of the Continental Shelf off the Indus Delta of Pakistan

Türk Denizcilik ve Deniz Bilimleri Dergisi

Cilt: 4 Sayı: 1 (2018) 33-45

Athar Ali KHAN^{1,*}, Rabea A. HAREDY¹

¹King Abdulazaiz University, Faculty of Marine Sciences, Department of Marine Geology, Jeddah, 21589, Saudi Arabia

ABSTRACT

Surface sediments collected from the continental shelf off the Indus delta. Pakistan, were studied to determine the organic carbon content. The results suggest that textural characteristics influence the distribution of the organic carbon content in the sediments of the Indus shelf. The surface sediments of Indus the shelf are predominantly silt, silty clay and clayey silty sand. Generally, inner shelf sediments contain 50-70% silt and ~20% clay. A belt of coarser sediments containing >50% sand occurs in the North West outer shelf region. The fine clay content in the sediments increases with depth, from 20% at the Indus

river mouth to 60% at the continental slope region. Sediments of the Indus shelf contain a relatively high (~1.0%) content of organic carbon, and their distribution indicates textural influence. The coarse. siltv sediments (50-70% silt) have a low organic carbon content of <1.0%. The fine-grained clayey sediments (20-40%) clay) are characterized by relatively high (1-2%) organic carbon content. The highest organic carbon content of 2-3% occurs in finegrained sediments of the outer shelf and the continental slope area, where 60% of the sediments are clayey sediments.

Keywords: Organic Carbon, Surface Sediments, Texture, Indus Continental Shelf

Article Info Received: 08 January 2018 Revised: 05 June 2018 Accepted: 08 June 2018

^{*}(corresponding author) *E-mail:* aamkhan@kau.edu.sa

1. INTRODUCTION

In continental shelf areas, sediments contain various amounts of organic carbon. The organic carbon content of these sediments is controlled by factors such as the primary productivity of the area (Calvert, 1987; Pederson and Calvert, 1990; Calvert et al., 1991), the bottom water dissolved oxygen (Demaison and Moore, 1980; Canfield, 1994), the sediment grain size (Premuzic et al., 1982; Mayer et al., 1985; Calvert, 1987; Ergin et al., 1993), the sedimentation rate (Betts and Holland, 1991) and sorption to mineral surfaces (Mayer, 1994; Keil et al., 1994a; Keil et al., 1994b). The relative importance of the different factors controlling organic matter accumulation is а controversial subject. Extensive studies on sedimentary organic matter have been undertaken to understand these factors (Hedges and Keil, 1995; Arthur et al., 1998; Hedges et al., 1999; Keil and Cowie, 1999; Vanderwiels et al., 2009; Zonneveld et al., 2010). Monsoonal upwelling and high primary productivity (Prell and Kutzbach, 1992; Qasim, 1977) characterize the study area in the northern Arabian Sea. Water columns at depths between 150 and 1,500 m have a broad, intense oxygen minimum zone (Wyrtki, 1971; Von Stackleberg, 1972). The occurrence of sediments rich in organic matter on the continental margins of Oman, Pakistan and India has been reported (Von stackelberg, 1972; Prell et al., 1992; Paropkari et al., 1992, 1993; Calvert et al., 1995). The organic carbon distribution in the surface sediments of the Arabian Sea and the controlling factors have been studied previously (Slater and Kropnick, 1984; Fontugne and Duplessy, 1986; Shimield et al., 1990; Calvert et al., 1995; Khan, 1999; Keil and Cowie, 1999; Cowie et al., 1999; Schulte et al., 2000; Cowie et al., 2009). Detailed studies of the biogeochemical characteristics of sediments undertaken by Cowie et al. (1999) confirmed the

accumulation of rich sediments of organic matter, predominantly of marine origin.

Earlier studies showed that surface productivity and dissolved oxygen are not the only factors controlling organic carbon distribution in this area. Paropkari et al. (1992) suggested considering other factors such as the texture of sediments. sedimentation rates, shelf-slope morphology, hvdrodvnamic clav mineralogy and processes. In the deep sea area of the Arabian Sea, Khan (1999) showed the temporal changes in organic carbon accumulation are coupled both to paleo productivity and sedimentation rates. Luckage et al. (2012) investigated a sediment core of the Indus Canyon and suggested that the Indus River affects productivity and therefore organic matter preservation in the coastal northeastern Arabian Sea. Here, surface sediments collected in the early eighties are re-evaluated with respect to organic carbon. The sediment texture and the carbonate content are published elsewhere (Khan, 1993). Textural data (Khan, 1993) is incorporated in this study to understand the distribution of the organic carbon content of the Indus shelf sediment. In the present study, the spatial distribution of organic carbon in surface sediments from offshore of the Indus delta is investigated and discussed based on sediment characteristics only. This study supports the hypothesis that sediment texture influences the organic carbon distribution in the surface sediments of the continental shelf off the Indus delta in Pakistan.

1.1. Regional Geological Setting

The northeastern part of the Arabian Sea between the Owen-Murray Ridge and the Indian continental margin is a passive continental margin. This margin is characterized by a 100-km-wide continental shelf and a stable, gently inclined continental slope. The shelf off the Indus delta is relatively flat. The shelf break occurs at the

135 m isobath (Giosan et al., 2006). The most pronounced morphological feature of the Indus shelf is the Indus Canyon (Islam, 1959; Shepard and Dill, 1966; Giosan et al., 2006). The Indus canyon crosses the ~100-km-wide shelf and the continental slope and extends as a fan-valley or channel-levee system onto the Indus fan. The Indus Canyon has played a significant role in funneling sediment from the shelf edge to deeper areas towards the Indus submarine fan (Islam, 1959; Nair et al., 1982; Wells and Coleman, 1984). The Indus River has been the major fluvial source of Arabian Sea sediments since the Early Miocene (Kolla et al., 1981; Clift et al., 2001; Clift et al., 2004; Garzanti et al., 2005).

2. MATERIAL AND METHOD

2.1. Study Area

The Indus shelf extends between Cape Monz and Sir Creek along the eastern continental margin of Pakistan as shown in Fig. 1. Prior to upstream damming, the Indus River supplied 450×10^6 t/year of sediments to the sea (Milliman et al., 1984; Milliman and Syvitski, 1992). The sedimentation rate on the Indus shelf is approximately 10 cm/year (Gibbs, 1981). Sediments on the modern Indus shelf are dominated by high silt and low carbonate content (Nair et al., 1982; Khan et al., 1993). The texture of the surface sediments divides the study area broadly into (i) near shore and inner shelf silt dominant sediments (~20-70 m) (ii) northwestern outer shelf zone and shelf break with sandy sediments (~70-130 m) and (iii) a deeper zone of continental slope featuring fine clayey sediments (Khan et al., 1993). Seasonal monsoonal upwelling results in high primary productivity and creates an oxygen minimum zone between depths of 200 to 1,500 m (Wyrtki, 1971; Von Stackleberg, 1972; Prell and Kutzbach, 1992).



Fig.1. Location map of sediment sample stations

2.2. Sampling and Analytical Methods

2.2.1. Sampling marine sediments

Surface sediment samples from the continental shelf, off the Indus delta, Pakistan (lat. 22° 47′ 34″ N to 24° 43′ 95″ N and long. 65° 58' 32" E to 67° 54' 0" E) were collected using a Peterson grab. The location of the sediment samples is given in Fig. 1. At stations 9, 23, 36, 39 and 40, owing to failure of a grab sampler, sediments were not collected. The grab samples were transferred into plastic jars and transported onshore. The top 4-5 cm of the sediment sample was scraped with a spoon and stored in a jar for analysis.

2.2.2. Textural Analysis

Grain size analysis was performed using standard grain size analysis methods (Folk 1980). Sand-sized (2 mm to 63 μ m), silt (63 to 4 μ m) and clay (<4 μ m) fractions were separated by wet sieving. A dispersant was added to 50 gm of the sample and wet sieved through a 63- μ m sieve. The fraction retained on the sieve was transferred to a pre-weighed evaporating dish. After drying, we calculated the coarser fraction (>63 μ m) in terms of sand content expressed as a percentage. Fine-sized particles (silt and clay) were characterized using their settling rates. Silt-sized material (<63 μ m) was washed through the sieve and collected in a litter cylinder. The fine material (<63 μ m) was concentrated by allowing a settling time and, after drying, was measured to determine the silt (<63 μ m) percentage. The remaining suspended material in the liquid was measured to determine the clay content after drying the decanted water through evaporation.

2.2.3. Organic Carbon Analysis

Sediment samples were dried at $60-70^{\circ}$ C. Bulk samples were ground to powder and homogenized. The powder sample was used for the analysis of organic carbon. A wet oxidation method using potassium dichromate and sulfuric acid as the oxidant was used (El-Wakeel and Riley, 1957). Gaudette *et al.* (1974) pointed out 100% yield can be achieved using this method.

3. RESULTS

3.1. Sediments

The surface sediments are gray, light gray, greenish gray, grayish green and brown in color. Sediment type and grain size characteristics of samples from the study area published earlier (Khan *et al.*, 1993) are given in Table 1. In the present study, the textural data have been studied to determine the relationship with organic carbon content. The texture of the sediments included coarse to fine sand, silt clay and silty clay. The surface sediments of the Indus shelf are dominated by the silt size (<63 μ m) fraction, which ranges between 40 and 80%. Texturally, sediments are classified as clayey silt (Folk, 1980), and the results are presented as a ternary diagram

(Fig. 2). Skeletal remains and broken pieces of foraminiferal ooze, pelecypods and gastropod shells are the predominant biogenic constituents. Sediment texture and the percentage of sand, silt and clay at each station are shown in Fig. 3, and their spatial distribution pattern is shown in Fig. 4. Based on sediment texture, the Indus shelf can be classified into three distinct sedimentary zones; (i) the innermost shelf zone (20-70 m), dominated by clayey silt (Fig. 4a), (ii) the northwestern shelf zone (70 - 100)m). dominated by coarse grained clayey sand and sand, (Fig. 4b) and (iii) the outer shelf zone (>100 m) and inner continental slope area, which contains silty clay sediments (Fig. 4c). Sediments from the stations proximal to the contour of 100 m depth in the eastern part of the study area in Zone I are clayey silt (~40% clay content, Fig. 4c) with some shell fragments and abundant mica flakes. The silt content is highest (~70%) in near-shore stations (Fig. 4a). This amount decreases to 30% towards the outer shelf stations, where the sand content increases and represents Zone II. Sediments on the western side of the Indus Canyon in Zone II are coarser and form a sand belt between the 50-100 m isobaths (Fig. 4b). In Zone II, the shelf sediments change from silt to fine sand with abundant skeletal material. Sediments collected from station numbers 18,19, 24, 26, 27, 28, 34, 37 and 38 show high sand content ranging from 50 to 70% (Fig. 3). The highest sand content, (~60%).occurs in sediments from stations 18. 24, 27, 28, 34 and 37 (Fig. 3). Sediments of Zone III in and around the lower reaches of Indus canyon at the continental slope are fine silt and clay with less shell material (Fig. 4c). The clay size fraction is dominant in sediments from stations 5, 6, 7, 10, 12, 16, 22, 25, 31, 32, 33 and 35 and ranges between 40 and 70% (Fig. 3).



Fig. 2. Textural Characteristics of the Indus shelf sediments



Fig. 3. Bar graph showing sand %, silt % and clay % in sediments of the Indus shelf



Fig. 4. Spatial distribution of sand, silt and clay % in the Indus shelf sediments

3.2. Organic Carbon Content

Sediment descriptions and organic carbon content are given in Table 1. The organic carbon content of the surface sediments from different stations is shown in a bar diagram (Fig. 5), and the surface distribution across the shelf is displayed in Fig. 6. The organic carbon content in the sediments from the continental shelf off the Indus delta ranges between 0.5% and 3.2%. Sediments present between the near shore zone and down to approximately 50 m water depth are low in organic carbon, which ranges from 0.5% to 1.0%. In general, organic carbon content in sediments increases from the inner shelf to the outer shelf regions. The sediments from the outer shelf stations (depth of ~100 m) exhibit comparatively high amounts of organic carbon (1-2%) compared to the inner shelf sediments, which exhibit an organic carbon content of 0.5% to 1.0%. Sediments from stations 5, 6, 7, 15, 16, 17, 25, 26, 34, 35, and 37 (Fig. 5) exhibit a high organic carbon content that ranges between >1.5% and 3.0%. The highest organic carbon content of ~2.50% occurs in sediments from stations 5, 6, 25, 26, 34, and 36 at depths between 100 and 200 m. Station 5 and 6 are located in the eastern part of the study area around the Indus canyon, whereas stations 25, 26, 34 and 36 are located westward and are clayey in nature.



Fig. 5. Organic carbon % and clay % in sediments of the Indus shelf



Fig. 6. Spatial distribution of sand, silt, clay and organic carbon % in the Indus shelf sediments

Station	Depth(m)	Latitude	Longitude	Types of	Color	*Sand	*Silt	*Clay	Organic
No		(N)	(E)	Sediments		(%)	(%)	(%)	Carbon
									%
1	13	23°.34′	67°.54′	Clay-silt	Dark grey	7.16	77.96	14.88	0.88
2	25	23°.23′	67°.43′	Clay-silt	Dark grey	11.05	77.26	10.42	0.63
3	26	23°.10′	67°.32′	Clay-silt	Dark grey	15.46	77.76	9.53	0.59
4	51	22°.56′	67°.22′	Clay-silt	Grey	9.16	73.74	4.3	0.8
5	113	22°.47′	67°.13′	Clay-silt	Grey	12.06	43.16	44.6	3
6	270	22°.54′	66°.58′	Silty-clay	Dark grey	2.88	19	78.12	2.5
7	112	23°.05′	67.11′	Silty-clay	Dark grey	1.65	36.78	62.13	2.4
8	49	23°.17′	67°.2′	Clay-silt	Dark grey	1.72	70	29	0.8
10	26	23°.41′	67°.37′	Clay-silt	Grey	1.85	49.23	48.92	0.63
11	295	23°.35′	67°.24′	Clay-silt	Light grey	4.2	64.4	30.6	0.8
12	712	23°.18′	67°.14′	Clay-silt	Brownish	1.2	60.1	38.4	0.9
					grey				
13	25	23°.45′	67°.18′	Clay-silt	Dark grey	1	65	33	0.95
14	80	23°.33′	67°.19′	Clay-silt	Dark grey	8	56	34	0.98
15	96	23°.22′	67°.01′	Clay-silt	Brownish	8	50	32	1.34
					grey				
16	137	23°.07′	66°.52′	Clay-silt	Brownish	9.72	49.18	41.38	1.5
					grey				
17	126	23°.25′	66°.39′	Clay-silt	Brownish	22.8	43.13	34.06	1.7
4.0		000 0 44	000 5/		grey		10 -		
18	90	23°.34′	66°.5′	Clay-sand	Brownish	57.7	18.7	23.44	0.88
10	50	000 441	078.01	Classicated	grey	50	40.04	20.44	0.75
19	59	23.44	07.0	Clay-sand	Brownish	50	18.21	30.11	0.75
20	23	23° 57'	67° 00'	Clay-silt	Groonish	0.3	70.2	10.0	0.5
20	23	23.31	07.09	Ciay-Sill	Greenish	9.5	10.2	10.0	0.5
21	30	24°.17′	67°.04′	Sandy-silt	Grev	12	75.5	4.5	0.5
22	22	24°.05′	66°.54′	Silty-clay	Grev	2.1	46.8	50.2	1.06
24	104	23° 42'	66° 32'	Silty-sand	Brownish	64.3	32 19	3.5	0.97
			00.01		arev	00	00	0.0	0.01
25	222	23°.30′	66°.25′	Sandy-	Greenish	18.5	7.4	73.3	2.6
				clay	grey				
26	216	23°.44′	66°.15′	Silty-sand	Grey	46.23	45.19	9.42	1.9
27	102	23°.55′	66°.24′	Silty-sand	Grey	63.11	33.34	6	1.2
28	80	24°.06′	66°.34′	Clay-sand	Greenish	70.04	11.19	18.6	0.75
					grey				
29	79	24°.19′	66°.44′	Sandy-silt	Grey	17.18	78.7	4.02	1
30	30	24°.29′	66°.65′	Clay-silt	Brown	2	70.69	27.28	0.5
31	27	24°.44′	66°.48′	Silty-clay	Grey	2.4	44.48	49.88	0.95
32	60	24°.32′	66°.37′	Silty-clay	Grey	1.2	46.44	50.96	0.82
33	60	24°.21′	66°.28′	Silty-clay	Grey	3.57	49.61	46.14	0.82
34	93	24°.09′	66°.17′	Clay-sand	Grey	42.5	27.4	30.1	2
35	174	23°.58′	65°.51′	Silty-clay	Brownish	19	30.4	50.6	1.8
27	157	210 101	66° 50'	Siltycoord	Dark brown	71.04	17 1	11 7	15
30	87	24 .12 21° 22'	66° 03'	Silty-sand	Brown	73.16	23.95	2.94	1.0
50	07	27.22	00.00	Jing-Sanu	BIOWII	10.10	20.00	2.04	

 Table 1. Location, textural characteristics and organic carbon content of the Indus Shelf sediments

*Published data (Khan, 1993)

4. DISCUSSION

Textural characteristics (sand silt and clay percentages) and the organic carbon content of the Indus shelf sediments from the study area are shown in bar diagrams in Figs. 3 and 5. The spatial distribution is also shown in Fig. 6. From the sand, silt, and clay percentages and their organic carbon content distribution, it is evident that sediments with a high clay percentage contain relatively high amounts of organic carbon. The correlation between the organic carbon content and the sand, silt, and clay percentages is shown in Fig. 7. These results suggest that organic carbon tends to be associated with fine sediments. Exceptions occurred only at station numbers 26 and 34 with ~2% organic carbon in coarser sediments (~50% sand). Station no. 26 is at a depth of 216 m, and station no. 34 is at 93 m. Increased organic carbon in the coarse sediments of these stations could be due to the bottom water and organic productivity, oxygen respectively.

It is known that particle size affects the organic carbon content of sediments (Trask, 1953; Brodovskiy, 1965; Van Andel, 1964; Hunt, 1996) and that clay contains more organic matter than coarse-grained sediments (Bush and Keller, 1981). In general, coarse sediments are expected to have lower organic matter content, either through less input or by winnowing of the organic content with fineparticles by erosive currents. grained Researchers (Listizen, 1972; Krissek and Scheidegger, 1983) have established a relationship between the clay content and the organic carbon content in marine sediments. Surface sediments and their organic carbon content determined in the present study (Fig. 6) reveal that bathymetry of the area and the sediment texture are the important controlling factors for the organic carbon content of sediments of the continental shelf off the Indus delta. The sediments are predominantly terrigenous. In the western part of the shelf, the sediments are coarser (silty sand) and

show increased biogenic content largely composed of skeletal material of calcareous organisms (Khan, 1993). The spatial distribution (Fig. 6) of sand, silt, clay and organic carbon and the R^2 values (Fig. 7) obtained for clay and sand suggest that the organic carbon content in sediments has more affinity towards fine clayey sediments. Sediments at station no. 6, 7 and 25 that are rich in clay content (50-70%) also show the highest organic carbon content value, i.e.,

>2.5%. There are stations (e.g., 34, 35, 37, and 38) from Zone II of the sandy belt that have relatively high organic carbon content. In contrast, the clayey sediments from stations 10, 30, 31, 32, 33 show relatively less organic carbon. Textural characteristics of the Indus shelf sediments suggest that there is a relationship between the grain size and the water depth. Sediments from shallow stations (~50 m water depth) are predominantly silty (50-70% silt) with low organic carbon ranging between 0.5 and 1%. Deeper stations from the outer shelf and the slope feature clay as the predominant sediment, along with higher organic carbon content (>2.0%). The textural dependency of the organic carbon content is evident from the fact that the finegrained sediments (<63 µm) have higher values (Fig. 5) compared to the coarser sediments. The relationship of a high concentration of organic carbon with high clay content is consistent with other published data (Van Andel, 1964; Reinson, 1975). Stackelberg (1972) and Paropkari et al. (1987) have studied the distribution of organic carbon in the sediments of the western shelf of India. According to Paropkari et al. (1987) the difference in the organic carbon observed between the inner shelf and outer shelf sediments could be attributed to the differences in the grain size of sediments, sedimentation rates, and activity of benthic population and to the role of oxygenated water. The high organic carbon content associated with fine-grained sediments of the slope region could be ascribed to the influence of the oxygen

minimum layer on the sea floor. Oxygen poor water (<0.5 ml/l) extends from 200 m to 1,500 m in the Arabian Sea (Slater and Kroopnick, 1984). This oxygen minimum layer prevents the destruction of organic carbon through oxidation. High values of organic carbon in clayey sediments found on the outer Indus shelf and slope region could be a result of both decreased bottom water oxygen and sediment texture.



Fig. 7. Correlation between sand, silt, clay and organic carbon %

5. CONCLUSION

The organic carbon content of the Indus shelf sediments suggests a close relationship with the sediment texture. The textural dependence of the organic carbon content is evident from the fact that the fine-grained sediments (<63 μ m) tend to have higher values when compared to the coarse sediments. The

organic carbon content in surface sediments of the outer shelf and the slope region is higher, and clay is the predominant sediment type. This study reveals the role of sediment texture in the distribution of the organic carbon content in sediments deposited off the Indus shelf of Pakistan. In this study, the observed spatial distribution pattern of the clay percentage in sediments suggests that the fine sediments are funneling through the Indus canyon and depositing deeper offshore. The organic carbon content distribution also indicates a similar process. The presence of fine, clayey sediments with a high OC content from the eastern stations in the canyon vicinity suggests that fine sediments may be acting as carriers and depositing.

ACKNOWLEDGEMENTS

First author would like to thank the ID4 cruise participants and their help in sample collection. We are grateful to DG NIO and acknowledge the help of M. Danish PSO, NIO Pakistan. Mr. Waqar and Mr. Jahangir of Geology Department Karachi University assisted and helped in the laboratory work for this research. Aid Zuberi and Talha A.Al-Dubai of Marine Geology Department King Abdul-Aziz University are thanked for using "SURFER" in producing the figures.

6. REFERENCES

Calvert, S.E., (1987). Oceanographic controls on the accumulation of organic matter in marine sediments In: "Marine Petroleum source Rocks". (J. Brooks & A.J. Fleet, eds), Geol. Soc. Spec. Publication, 26: 137-151.

Pedersen, T.F., Calvert, S.E., (1990). Anoxia vs productivity: what controls the formation of organicrich sediments and sedimentary rocks? *Am. Assoc. Petrol Geol. Bull.* 74: 454–466.

Calvert, S.E., Karlin, R.E., Toolin, L.J., Donahue, D.J., Southon, J.R., Vogel, J.S., (1991). Low organic carbon accumulation rates in Black Sea sediments. *Nature* 350 (6230): 692-695.

Demaison, G.J., Moore, G.T., (1980). Anoxic environments and oil source bed genesis. *Am. Assoc. Petrol Geol. Bull.* 64: 1179–1209.

Canfield, D.E., (1994). Factor influencing organic carbon preservation in marine sediments. *Chem. Geol.* 114: 315–329.

Premuzic, E.T., Benkovitz, C.M., Gaffney, J.S., Walsh, J.J., (1982). The nature and distribution of organic matter in the surface sediments of world oceans and seas. *Org. Geochem.* 4: 63–77.

Mayer, L. M., Rahim, P.T., Guerin, W., Macko, S.A., Walting, L., Anderson, F.E., (1985). Biological and granulometric controls on sedimentary organic matter of an intertidal mudflat. *Estuarine Coastal Shelf Science* 20: 491-503.

Ergin, M., Bodur, M.N., Ediger, D., Ediger, V., Yılmaz, A., (1993). Organic carbon distribution in the surface sediments of the Sea of Marmara and its control by the inflows from adjacent water masses. *Marine Chemistry* 41:311-326.

Betts, J.N., Holland, H.D., (1991). The oxygen content of ocean bottom waters the burial efficiency of organic carbon and the regulation of atmospheric oxygen. *Global and Planetary Change* 5: 5–18.

Mayer, L.M., (1994). Relationship between mineral surfaces and organic carbon concentrations in soils and sediments. *Chem. Geology* 114: 347-363.

Keil, R.G., Tsamakis, E., Fuh, C.B., Giddings, J.C., Hedges, J.I., (1994a). Mineralogical and textural controls on the organic composition of coastal marine sediments: hydrodynamic separation using SPLITT fractionation. *Geochimica et Cosmochimica Acta* 58: 879-893.

Keil, R.G., Montluion, D.B., Prahl, F.G., Hedges, J.I., (1994b). Sorptive preservation of labile organic matter in marine sediments. *Nature* 370: 549-552.

Hedges, J.I., Keil, R.G., (1995). Sedimentary organic matter preservation: An assessment and speculative synthesis. *Mar. Chem.* 49: 81–115.

Arthur, M.A, Dean, W.E., Laarkamp, K., (1998). Organic carbon accumulation and preservation in surface sediments on the Peru margin. *Chem. Geol.* 152: 273–286.

Hedges, J.I., Hu, F.S., Devol, A.H., Hartnett, H.E., Tsamakis, E., Keil, R.G., (1999). Sedimentary organic matter preservation: A test for selective oxic degradation. *Am. J. Sci.* 299: 529–555.

Keil, R.G., Cowie, G.L., (1999). Organic matter preservation through the oxygen- deficient zone of the NE Arabian Sea as discerned by organic carbon:

mineral surface area ratios. *Marine Geology* 161: 13–22.

Vanderwiele, S., Cowie, G.L., Soetaert, K., Middelburg, J.J., (2009). Amino acid biogeochemistry and organic matter degradation state across the Pakistan margin oxygen minimum zone. *Deep-Sea Res Pt II* 56(6–7): 376–392.

Zonneveld, K.A.F., Versteegh, G.J.M., Kasten, S., Eglinton, T.I., Emeis, K.C., Huguet, C., Koch, B.P., de Lange, G.J., de Leeuw, J.W., Middelburg, J.J., Mollenhauer, G., Prahl, F., Rethemeyer, J., Wakeham, S., (2010). Selective preservation of organic matter in marine environments; processes and impact on the fossil record. *Biogeosciences* 7: 1-29.

Prell, W.L., Kutzbach, J.E., (1992). Sensitivity of the Indian monsoon to forcing parameters and implications for its evolution. *Nature* 360: 647–652. doi: 101038/360647a0

Qasim, S.Z., (1977). Biological productivity of the Indian Ocean Indian. J. Mar. Sci. 6: 122-I37.

Wyrtki, K., (1971). Oceanographic Atlas of the International Indian Ocean Expedition National Science Foundation US Government Printing Office Washington DC, 531.

Von Stackelberg, U., (1972). Facies of sediments of the Indian-Pakistan continental margin (Arabian Sea). *Meteor Forsch Ergebnisse Reihe* 9: 1-73.

Paropkari, A.L., Babu, C.P., Mascarenhas, A., (1992). A critical evaluation of depositional parameters controlling the variability of organic carbon in Arabian Sea sediments. *Marine Geology* 107(3): 213-226.

Paropkari, A.L., Babu, C.P., Mascarenhas, A., (1993). New evidence for enhanced preservation of organic carbon in contact with oxygen minimum zone on the western continental slope of India. *Marine Geology* 111(1–2): 7-13.

Calvert, S.E., Pedersen, T.F., Naidu, P.D., von Stackelberg, U., (1995). On the organic carbon maximum on the continental slope of the eastern Arabian Sea. *Journal of Marine Research* 53: 269–296.

Slater, R.D., Kroopnick, P., 1984. Controls on dissolved oxygen distribution and organic carbon deposition in the Arabian Sea In: "Marine Geology and Oceanography of Arabian Sea and Coastal Pakistan", (J.D. Milliman, B.U. Haq, eds.), Nostrand Reinhold, New York. Fontugne, M.R., Duplessy, J.C., (1986). Variations of the monsoon regime during the upper Quaternary:" evidence from carbon isotopic record of organic matter in north Indian Ocean sediment cores. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* 56: 69-88.

Shimmield, G.B., Price, N.B., Pedersen, T.F. (1990). The influence of hydrography bathymetry and productivity on sediment type and composition of the Oman Margin and in the Northwest Arabian Sea. In: "The geology and tectonics of the Oman region", (A.H.F. Robertson, M.P. Searle, A.C. Ries, eds), Geological Society Special Publication No 49, pp. 759–769, London.

Khan, A.A., (1999). Accumulation of organic carbon in northwestern Arabian Sea sediments. *Pakistan Journal of Hydrocarbon Research* 11: 51-58. georefid: 2008-126645

Cowie, G.L., Calvert, S.E., Pedersen, T.F., Schulz, H., von Rad, U., (1999). Organic content and preservational controls in surficial shelf and slope sediments from the Arabian Sea Pakistan Margin. *Mar. Geol.* 161: 23–38.

Schulte, S., Mangelsdorf, K., RullkoÈtter, J., (2000). Organic matter preservation on the Pakistan continental margin as revealed by biomarker geochemistry. *Organic Geochemistry* 31(10): 1005-1022.

Cowie, G.L., Mowbray, S., Lewis, M., Matheson, H., McKenzie, R., (2009). Carbon and nitrogen elemental and stable isotopic compositions of surficial sediments from the Pakistan margin of the Arabian Sea. *Deep-Sea Research* 56: 271–282. doi:101016/jdsr2200805031

Lückge, A., Deplazes, G., Schulz, H., Scheeder, G., Suckow, A., Kasten, S., Haug, G.H., (2012). Impact of Indus River discharge on productivity and preservation of organic carbon in the Arabian Sea over the twentieth century. *Geology* 40(5): 399–402.

Khan, A.A., Memon, M.G., Danish, M, Inam, A., (1993). Distribution of surface sediments off Indus delta on the continental shelf of Pakistan. *Pakistan Journal of Marine Sciences* 2(1): 33–39.

Giosan, L., Constantinescu, S., Clift, P.D., Tabrez, A.R., Danish, M., Inam, A., (2006). Recent morphodynamics of the Indus delta shore and shelf. *Continental Shelf Research* 26: 1668–1684

Islam, S.R., (1959). The Indus submarine canyon. *Oriental Geography* 3: 101–104.

Shepard, F.P., Dill, R.F. (1966). *Submarine canyons and other sea valleys*. Rand McNally and Co, Chicago.

Nair, R.R., Hashimi, N.H., Rao, P.C., (1982). Distribution and dispersal of clay minerals on the western continental shelf of India. *Mar. Geol.* 50: 1-9.

Wells, J.T., Coleman, J.M., 1984. Deltaic morphology and sedimentology with special reference to the Indus River delta. In: "Marine Geology and Oceanography of Arabian Sea and Coastal Pakistan", (J.D. Milliman, B.U. Haq, eds.), pp. 85–100, Nostrand Reinhold, New York.

Kolla, V., Ray, P.K., Kostecki, J.A., (1981). Surficial sediments of the Arabian Sea. *Mar. Geol.* 41:183–204.

Clift, P.D., Shimizu, N., Layne, G., Blusztajn, J., (2001). Tracing patterns of unroofing in the Early Himalaya through microprobe Pb isotope analysis of detrital K-feldspars in the Indus Molasse. *Earth Planet Sci. Lett.* 188: 475–491.

Clift, P.D., Campbell, H., Malcolm, S.P., Andrew, C., Zhang, X., Hodges, K.V., Ali Athar, K., Charlotte, M.A., (2004). Thermochronology of the modern Indus River bedload; new insight into the control on the marine stratigraphic record. *Tectonics* 23. doi: 101029/2003TC001559

Garzanti, E., Vezzoli, G., Ando, S., Paparella, P., Clift, P.D., (2005). Petrology of Indus River sands: a key to interpret erosion history of the Western Himalayan Syntaxis. *Earth and Planetary Science Letters* 229: 287–302.

Milliman, J.D., Quraishee, G.S., Beg, M.A.A. (1984). Sediment discharge from the Indus River to the ocean: past present and future In: "Marine Geology and Oceanography of the Arabian Sea and Coastal Pakistan Van Nostrand Reinhold", (B. U. Haq, J. D. Milliman, eds), pp. 65-70, New York.

Milliman, J. D., Syvitski, P. M., (1992). Geomorphic/ tectonic control of sediment discharge to the ocean: the importance of small mountainous rivers. *Journal of Geology* 100: 525–544.

Gibbs, J.R., (1981). Sites of river derived sedimentation in the Ocean. *Geology* 9: 77-80.

Folk, R.L. (1980). *Petrology of Sedimentary rocks*. Hemphill Publishing Company, Austin, Texas.

El–Wakeel, S.K., Riley, J.P., (1957). The determination of organic carbon in marine muds. *J. Cons. Int. Explor. Mer.* 22(2): 180–183.

Gaudette, H., Muller, G., Stoffers, P., (1974). An inexpensive titration method for the determination of organic carbon in recent sediments. *J. Sed. Pet.* 44: 249-253

Trask, P.D., (1953). The sediments of the Western Gulf of Mexico. *Part II: Papers Phys Oceanogr and Meteorol Mass Inst Technol Woods Hole Oceanogr Inst* 12:51.

Brodovskiy, O.K., (1965). Accumulation of organic matter in bottom sediments. *Marine Geology* 3:33-82.

Van Andel, T.H., (1964). Recent marine sediments of the Gulf of California. In: "Marine geology of the Gulf of California", (T.H. Van Andel, G.D. Shor, eds.), *Am. Assoc. Petrol Geol. Mem.* 3:216.

Hunt, J.M. (1996). *Petroleum Geochemistry and Geology*, Freeman Press, New York.

Bush, W.H., Keller, G.H., (1981). The physical properties of Peru-Chile continental margin sedimentsthe influence of coastal upwelling on sediment physical properties. *Jour. Sedimentary Petrology* 51:705-719.

Lisitzin, A.P., (1972). Sedimentation in the world Ocean. Soc. Economic Paleontologists and Mineralogists Special Publication 17: 218.

Krissek, L.A., Scheidegger, K.F., (1983). Environmental controls on sediment texture and composition in low oxygen zones off Peru and Oregon In: "Coastal Upwelling: Its Sediment Record Part B: Sediment Records of Ancient Coastal Upwelling Plenum", (J. Thiede, E. Suess, eds.), pp. 163-180, New York.

Reinson, G. E., (1975). Geochemistry of muds from shallow restricted estuary Australia. *Mar. Geol.* 19: 297-314.

Paropkari, A.L., Rao, C.M., Murty, P.S.N. (1987) Environmental controls on the distribution of organic matter in recent sediments of the western continental margin of India. In: "Petroleum geochemistry and exploration in the Afro-Asian region", (R.K. Kumar, P. Dwivedi, V. Banerjee & V. Gupta, eds), pp. 347–361, Rotterdam.