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A Comparative Study on Maximum Shear Strain Distribution in a Layered Viscoelastic Medium

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Abstract - For practical calculations, the ovaling and racking deformations on buried structures are estimated by estimating the maximum shear strain (γ_{max}) in imperforated ground using vertically incident shear waves (V_{s}). The relationship between maximum shear strain (γ_{max}) and depth in a uniformly elastic half-space is a function of shear-wave velocity in the uniform medium and acceleration history in free field. In this study, the relationship between the distribution of γ_{max} in elastic half space and in layered medium is examined by supposing that the travel time of shear waves from the free boundary to a depth of interest in a layered medium is equal to d/V_s . Different variety of strong ground motion time histories and three V_s profiles are used for analyses. The effect of nonlinear material response to shearing was investigated by using the method of equivalent linearization.

Keywords: Shear strain, underground structures, ovaling, racking, free field.

1. Introduction

Underground structures serve as means of transportation, sanitation, irritation and storage. Recent earthquakes such as Kobe (1995), Kocaeli (1999) and Chi-Chi (1999) showed the vulnerabilities of these structures to severe seismic loads. Ovaling and racking deformations due to vertically induced shear waves are the main causes of seismic damages (Hashash et.al, 2001). Analytical solutions using free field approach and soil-structure interaction have been proposed to calculate ground deformations. In free-field approach, it is assumed that the structure conforms to the surrounding ground deformations. Due to lack of in-situ ground deformation data, simplified formulas are used to calculate γ_{max} . γ_{max} is generally related to the formula of Newmark, supposing unidirectional wave propagation in an unbounded uniformly elastic medium, such that

$$\gamma_{\max} = v_{\max} / C \tag{1}$$

where v_{max} is the maximum particle velocity, and *C* is the apparent wave velocity (Newmark, 1968; St.John and Zahrah, 1987; Hashash et.al, 2001). For ovaling analyses, a reasonable simplification for computation of γ_{max} is to assume vertically incident shear waves propagating through horizontal layers (Wang, 1993; Hashash et.al, 2001). In that case, *C* is approximately equal to V_s , the propagation velocity of shear waves in a layer. However, equation (1) yields conservative estimates for γ_{max} , especially in shallow ranges of depth. The ground γ_{max} spectrum provides the

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relationship between γ_{max} and depth for an acceleration record on uniformly elastic half space (Chen and Hou, 1992). By defining the seismic hazard in terms of shear-strain spectrum, it is possible to estimate γ_{max} reasonably for any given depth in a uniformly elastic half space. In order to estimate shear strain in actual sites which generally consists of layered geological formations, a relationship relating γ_{max} spectra in uniformly elastic half space to the distribution of γ_{max} in layered soil profiles is necessary. A reasonable method may be the use of equivalent travel-time concept proposed by Imai et al. (1981) for the estimation of maximum shear-stress in layered medium. This is examined for estimation of γ_{max} in layered soil profiles. Furthermore, the effect of nonlinearity of soil behavior on the accuracy of this method is investigated.

2. Shear-Strain Spectrum

The equation of particle motion due to vertically incident shear waves that propagate with velocity V_s in a linearly elastic medium can be written as

$$\frac{\partial^2 u}{\partial t^2} = V_s^2 \frac{\partial^2 u}{\partial y^2} \tag{2}$$

where, u(t,y) depicts the horizontal displacement of a particle located at coordinate y and at time t due to the disturbance propagating in the vertical direction. The relationship between shear strain (γ) at any depth (d) and shear wave velocity is

$$V_{s} \cdot \gamma^{F}(\omega, d/V_{s}) = \frac{a^{F}(\omega)}{\omega} \cdot \sin\left(\frac{\omega}{V_{s}}d\right)$$
(3)

In equation (3), a^{F} denotes the complex Fourier coefficient of acceleration history (*a*) in free field. Equation (3) can be expressed in time domain by

$$V_{s} \cdot \gamma(t, d/V_{s}) = \frac{1}{2} [v(t + d/V_{s}) - v(t - d/V_{s})]$$
(4)

where v(t) is the velocity history in free field. Hence, γ_{max} is a function of d/V_s , the travel time of vertically incident shear waves from free boundary to a point at depth d (Chen and Hou, 1992).

3. The Concept of Equivalency in Travel Time

Equations (3) and (4) are applicable for uniformly elastic half space. According to Imai et al. (1981), the maximum shear stress (τ_{max}) at any given depth in layered soils can be estimated by supposing that the travel time of vertically incident shear waves from the free boundary to the depth of interest is equivalent to the travel time in a hypothetically uniform elastic medium. This travel time is calculated by the formula,

$$\frac{d}{V_s} = \sum_{i=1}^l \frac{d_i}{V_i} \tag{5}$$

for the estimation of τ_{max} in the *l*th layer below the ground surface. Then, γ_{max} can be calculated by

$$\gamma_{max} = \frac{\tau_{max}}{\rho_l \times (V_s)_l^2} \tag{6}$$

where ρ_1 and $(V_s)_I$ are respectively the density and the shear-wave velocity of the layer of interest. Hence, d/V_s calculated by equation (5) is first substituted in equations (3) or (4) for the estimation of $V_s\gamma(t)$. Then, by supposing that $V_s = (V_s)_I$, the maximum absolute value of $V_s\gamma(t)$ yields γ_{max} for the l^{th} layer. The procedure can be simplified if the relationship between d/V_s and $V_s\gamma_{max}$ is first computed for any given a(t), and then d/V_s calculated by equation (5), is used for interpolation in the calculated range of $V_s\gamma_{max}$.

Equations (3) to (5) are based on the assumption of linearly elastic medium whereas the geological formations are expected to show nonlinear stress-strain response to severe shearing during seismic events. The effect of nonlinear material behavior on γ_{max} can be involved in calculations by the well-known method of equivalent linearization (Schnabel et al., 1972). The accuracy of the method was investigated by the comparisons with the 1-dimensional siteresponse analyses. The program PROSHAKE, computing the dynamic site response by the method of equivalent linearization (EduPro Civil Systems, 1998), was used for the determination of γ_{max} in each layer. In the analyses, the ground motion was defined for a hypothetical station on the ground surface, so that $\gamma(t)$ was computed by the deconvolution technique. The strain-dependent degradation in shear modulus (G), and the increase in material damping (ξ) were modeled by the empirical relationship of Vucetic and Dobry (1991), applicable to lightly overconsolidated fine soils with a Plasticity Index of 15, so that the strain-dependent reduction in material rigidity is significant in the dynamic response analyses. The representative value for ξ and the reduction in G are estimated by supposing that the cyclic amplitude of γ is equal to 65% of γ_{max} . $(V_s)_l$ is proportional to the square-root of G in the equivalent-linear method of analysis due to the relation $G = \rho \times (V_s)$. The tolerable relative error, the limit for relative difference between the results of successive iterations, was accepted as 1%. In these analyses, the secant shear modulus is changed for each layer until their values are consistent with γ_{max} induced in each layer. The same iterative procedure was implemented in a spread sheet program for the calculation of γ_{max} by the shear-strain spectrum.

4. Analyses

A number of accelerograms were selected from the strong-motion databases PEER and DAPHNE to collect a sample for the empirical determination of parameters. The emphasis was placed on using the records that involve significant ground-motion amplitudes for buried structures. The PEER database provided filtered records. The records are attributed to the seismic events presented in Table 1. The sites of records were classified according to the site classification system of *International Building Code* (ICC, 2009). This widely known system is related to the parameter *Vs30*, the weighted harmonic mean of V_s to the depth of 30 m (Borcherdt, 1994). The site class D is pertinent to stiff-soil (180 m/s $\leq Vs30 \leq 360$ m/s) condition. The frequencies above 15 Hz can be heavily contaminated for accelerograms by high-frequency surface waves and by waves reflected from stratified formations. The highfrequency contamination yields amplitudes that are not consistent with the assumptions about the properties of viscoelastic media considered in analyses, and can be suppressed by a low-pass filter to improve the coherence between the velocity profiles and the accelerogram defining the free-field motion (Silva, 1988). Therefore, the records were filtered by a 4-pole/4-pole low-pass Butterworth filter (Akkar and Bommer, 2007). The ground motion records presented in Table 1 were supposed to be the records of ground motion in free field after suppressing the frequencies above 15 Hz. Three real velocity profiles, shown in Figure 1, were considered in site-response analyses. The first profile (Figure 1.a) is adopted from the site *Bell-La Bulk Mail* presented by Nigbor and Swift (2001). The second and third profiles (1.b, and 1.c) are adopted from the sites *Gilroy* #2 (USGS) and Palo Alto Veterans Hospital presented by Boore et al. (2003). An elastic half space representing the soft rock was considered to be lying beneath the viscoelastic layers. The distribution of V_s with respect to d/V_s in three sites is shown in Figure 2. d was taken as the average depth of each uniformly viscoelastic layer. The material density of all actual formations was set to 2000 kg/m³, so that the results of analyses can be solely attributed to the distribution of V_s .

The records 6 and 16 (Table 1) were scaled by the factor 0.60 to keep γ_{max} below 1% because the accuracy of equivalent linearization in site-response analysis can be considerably low if γ_{max} is greater than 1% (Gerolymos and Gazetas, 2005). Due to the size limitations in the computer program, the record numbered as 20 in Table 1 was cropped to shorten the record length by the 70 seconds that involve the most significant amplitudes of ground motion. The cropped record was filtered at high-pass frequency of 0.1 Hz.

Rec. No.	Event No.	Station	Vs30	SC	Event Name	Event Date	Ms	Depth (km)
1	1	Gormon - Oso Pump Plant	308	D	San Fernando	9.02.1971	6.6	13.0
2	2	Capitola	288	D	Loma Prieta	18.10.1989	7.1	17.5
3	2	Gilroy Array #2	270	D	Loma Prieta	18.10.1989	7.1	17.5
4	2	Gilroy Array #3	349	D	Loma Prieta	18.10.1989	7.1	17.5
5	2	Gilroy - Historic Bldg.	338	D	Loma Prieta	18.10.1989	7.1	17.5
6	3	Eureka - Myrtle & West	338	D	Cape Mendocino	25.04.1992	7.1	9.6
7	4	Amboy	271	D	Landers	28.06.1992	7.4	7.0
8	4	Fort Irwin	345	D	Landers	28.06.1992	7.4	7.0
9	4	Hemet Fire Station	338	D	Landers	28.06.1992	7.4	7.0
10	5	Inglewood - Union Oil	316	D	Northridge	17.01.1994	6.7	17.5
11	5	LA - Century City CC North	278	D	Northridge	17.01.1994	6.7	17.5
12	5	LA - E Vernon Ave	308	D	Northridge	17.01.1994	6.7	17.5
13	6	KJMA	312	D	Kobe	16.01.1995	6.9	17.9
14	6	Kakogawa	312	D	Kobe	16.01.1995	6.9	17.9
15	6	Takarazuka	312	D	Kobe	16.01.1995	6.9	17.9
16	6	Takatori	256	D	Kobe	16.01.1995	6.9	17.9
17	6	Shin-Osaka	256	D	Kobe	16.01.1995	6.9	17.9
18	7	1612	197	D	Kocaeli	17.08.1999	7.7	15.9
19	7	5903	325	D	Kocaeli	17.08.1999	7.7	15.9
20	8	СНУ036	233	D	Chi-chi	20.09.1999	7.6	6.8

SC: site class (International Code Council).

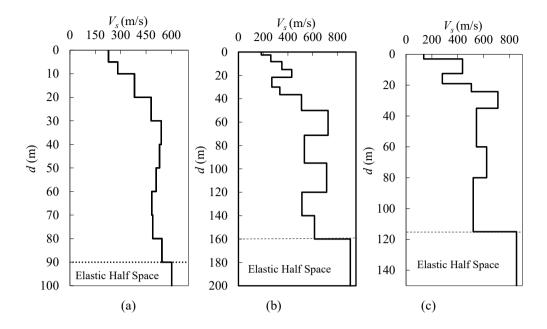


Figure 1: Velocity profile used for the site response analyses for (a) *Bell - La Bulk Mail* site, for (b) *Gilroy* #2 (USGS), and for (c) *Palo Alto Veterans Hospital*.

Figure 3 shows the range of geometric means of maximum strain ratio ($\gamma_{max-l}/\gamma_{max}$) for two horizontal components of ground motion used in PROSHAKE analyses, such that a linear relationship between stress and strain is considered. γ_{max-l} denotes maximum absolute shear strain computed by PROSHAKE. $V_s\gamma_{max}$ is calculated by equation (3) or (4). Then, γ_{max} is calculated by supposing that $V_s = (V_s)_l$. The comparisons of figures (2) and (3) reveal that this analysis is very sensitive to the variability in V_s . An investigation of the ratio between V_s in soft layers and that in stiffer layers shows that the method may slightly underestimate γ_{max} in soft layers located beneath much stiffer layers. However, the predictions are conservative for most of the layers. For instance, the lowest V_s ratio (0.62) occurs at the travel time 0.07 s at the site *Gilroy* #2 (USGS), such that $\gamma_{max-l}/\gamma_{max}$ is almost 1.0. Whereas, at the site *Bell - La Bulk Mail site*, the V_s ratio is considerably low (0.64) for $d/V_s = 0.05$ s, which is close to the minimum ratio at the site *Gilroy* #2 (USGS), but the underestimation of γ_{max} is negligible. One important observation is that, the scattering of $\gamma_{max-l}/\gamma_{max}$ for any given d/V_s is very limited.

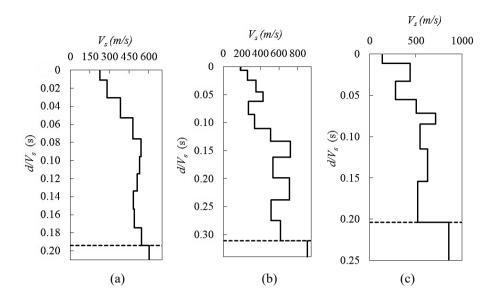


Figure 2: The travel times for middle geological layers on the sites (a) *Bell - La Bulk Mail*, (b) *Gilroy #2 (USGS)*, and (c) *Palo Alto Veterans Hospital*.

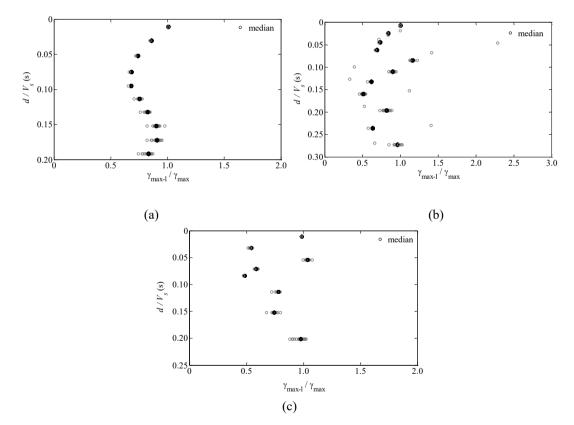


Figure 3: The ranges of $\gamma_{max-l}/\gamma_{max}$ computed by supposing linear response of layers on the sites (a) *Bell - La Bulk Mail*, (b) *Gilroy #2 (USGS)*, and (c) *Palo Alto Veterans Hospital*.

In order to investigate the effect of nonlinearity of soils' response to shearing, the same profiles were analyzed by using the relationships between ξ , (G/G_{max}) and cyclic amplitude of γ that were suggested by Vucetic and Dobry (1991). Figure 4 shows the scattering of $\gamma_{max-l}/\gamma_{max}$ by d/V_s . The comparisons of figures (3) and (4) reveals that the scattering of $\gamma_{max-l}/\gamma_{max}$ and the underestimation of equation (3) for embedded soft layers can be more prominent if the soils behave nonlinearly. This can be explained by the reduction in *G*, consequently in equivalent V_s . Except for these soft layers embedded in much stiffer layers, equation (3) usually yields a reasonable or conservative estimate for γ_{max} .

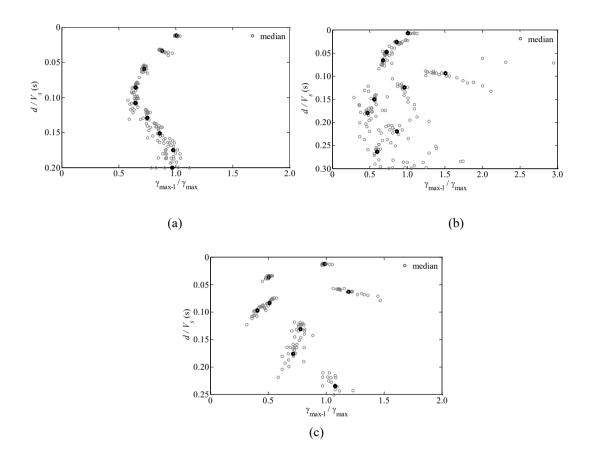


Figure 4: The ranges of $\gamma_{max-l}/\gamma_{max}$ computed by supposing nonlinear response of layers on the sites (a) *Bell - La Bulk Mail*, (b) *Gilroy #2 (USGS)*, and (c) *Palo Alto Veterans Hospital*.

6. Conclusions

The assumption that the parameter $V_s \gamma_{max}$ in any soil layer is strongly related to the travel time of shear waves from ground surface to the depth of interest yields a method for the estimation of γ_{max} , provided that the shear-strain spectrum suggested by Chen and Hou (1992) is available. Therefore, in calculation of seismic hazard for buried structures and in the selection of ground-motion records, the emphasis should be given to this spectrum. However, strain ratio is sensitive to the site layering, and may show limited scattering. The procedure explained in this study may result in the underestimation of γ_{max} in soft layers embedded in much stiffer layers. Detailed site-response analyses for the calculation of γ_{max} in such soft layers are observed to be crucial. Otherwise, the method yields either reasonably accurate or conservative estimates for γ_{max} in layered geological formations. The accuracy of the method is very satisfactory for almost uniform V_s profiles. All conclusions are based on the assumption that the ground motion is induced by a field of vertically propagating shear waves. Hence, the proposed method can be acceptable for first-order estimations of γ_{max} , particularly in the absence of the data necessary for more accurate analysis methods, if the seismic hazard is expressed by a shear-strain spectrum.

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